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## HIMALAYAN AND TRANS-HIMALAYAN GLACIER FLUCTUATIONS SINCE AD 1812

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### ABSTRACT

Historical records of the fluctuations of glaciers in the Himalayas and Trans-Himalayas date back to the early 19th century. Local and regional syntheses of 112 of these fluctuation records are presented in this study. The local syntheses deal with fluctuations of glaciers in Kanchenjunga-Everest, Garwhal, Lahaul-Spiti, Kolahoi, Nanga Parbat, Karakoram (north and south sides), Rakaposhi-Haramosh, Batura Mustagh, and Khunjerab-Ghujerab. Regional syntheses deal with the composite record and the differentiation of

records by glacier type (longitudinal versus transverse) and regional setting (Himalayan versus Trans-Himalayan). In a gross regional sense Himalayan and Trans-Himalayan glaciers have been in a general state of retreat since AD 1850. Filtering of the fluctuation records with respect to glacier type and regional setting reveals that the period AD 1870 to 1940 was characterized by alternations in the dominance of retreat, advance, and standstill regimes.

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### INTRODUCTION

The glaciers of high Asia comprise by area 50% of all glaciers outside of the polar regions, and they contain approximately 33 times the areal cover of the glaciers in the European Alps (Wissman, 1959). The vast glacier coverage included in two of the largest components of the high Asia glacier complex, the Himalayas and the Trans-Himalayas (Figure 1) is characterized by a regional firn limit that ranges in altitude from 5500-5600 m for Everest (Müller, 1958), to 5200-5700 m for the Garwhal (Grinlinton, 1914), to 4500-4700 m for Nanga Parbat (Finsterwalder, 1937), to 5000-5200 m for the south side of the Karakoram, and to 5500 m for the north side of the Karakoram (Visser and Visser-Hooft, 1938). Both the wide altitudinal range and the diverse orientations of the glaciers in

the Himalayas and the Trans-Himalayas (Figure 2) make these regions excellent sites for the study of glacier fluctuations.

Historical records of glacier fluctuations in the Himalayas and Trans-Himalayas extend back over 150 years. The earliest studies concerned with the movement of glacier termini were made for Chong Kumdan Glacier in AD 1812 (Izzet Ullah, 1842) and Milam Glacier (Hodgson, 1822). More modern references, cited by glacier in the Appendix, provide data on the fluctuation history of 112 glaciers. Several papers summarize the fluctuation record of subgroups of this total; notable amongst these are Mason (1930a), Visser and Visser-Hooft (1938), Mercer (1963), and Tewari (1971).

Mason (1930a), using historical records,

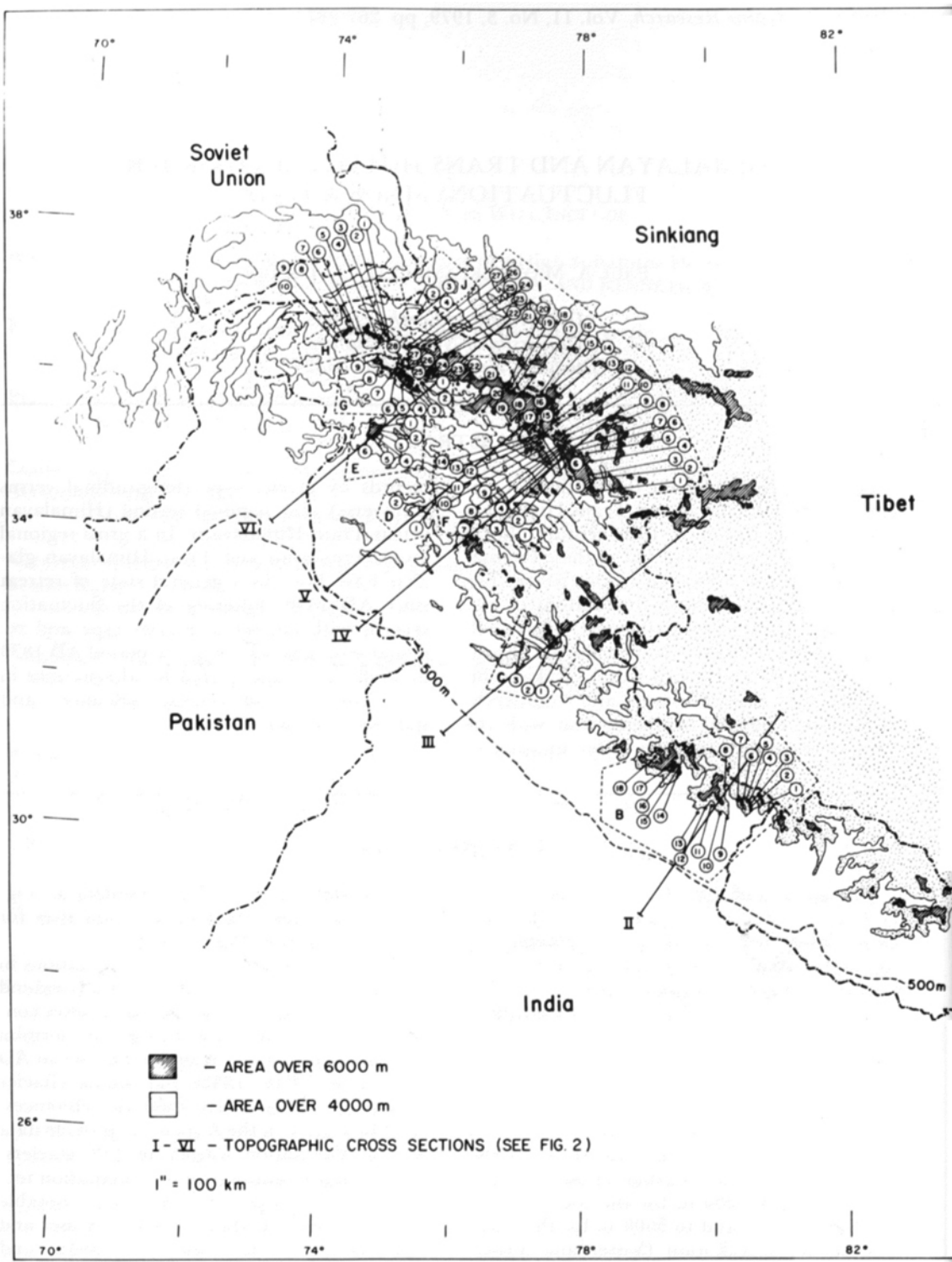
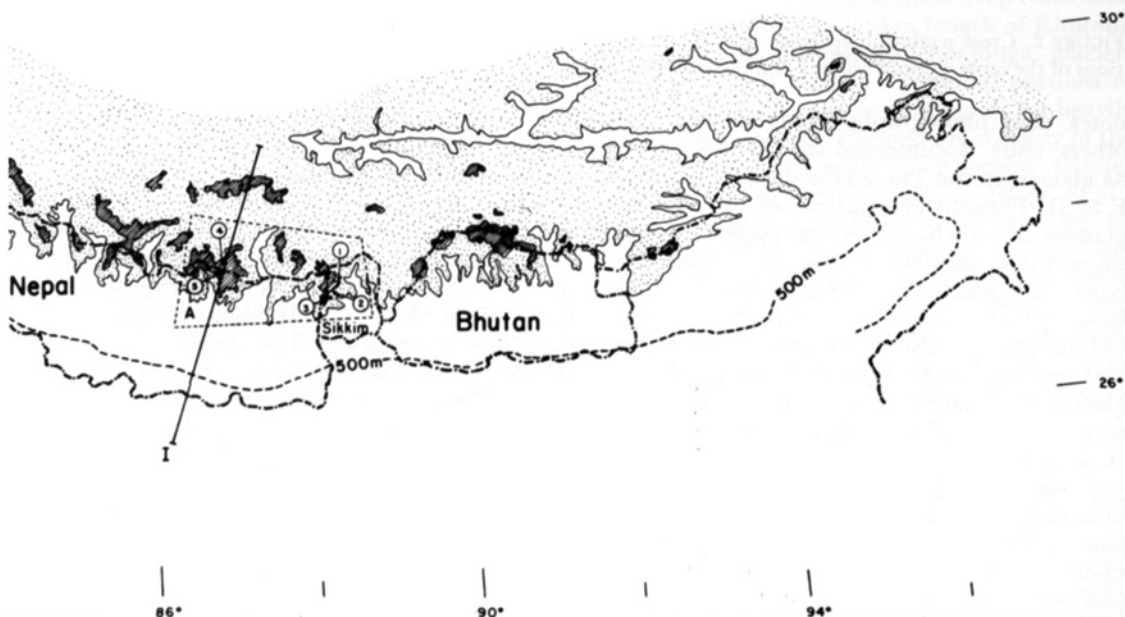


FIGURE 1. Glacier locations.

Himalaya		Trans - Himalaya	
Everest-Kanchenjunga - Region A		South Side Karakoram - Region F	
1. Zemu		1. South Saser	
2. Alukthang		2. North Saser	
3. Guicha		3. Taghman	
4. Jungpu		4. Manlung	
5. Khumbu		5. Namlung	
Garwhal - Region B		6. Changmolung	
1. Baling		7. Chogolala	
2. Sona		8. Syethang	
3. Naulphu		9. Mamostang	
4. Nipchungkang		10. Shosholing	
5. Kharsa		11. La Yoghma	
6. Chingchingmauri		12. South Terong	
7. Shankalpa		13. Sheikar Chorten	
8. Milan		14. North Terong	
9. Poting		15. Major tongue in Nuhra Valley	
10. Kaphni		16. Siachen	
11. Pindari		17. Rgyong Valley glaciers	
12. Burghal-Maktoli		18. Bilafond and Chumik	
13. Mrigithuni		19. Sherpingang	
14. Bhagirath-Kharak		20. Khondokoro	
15. Satopanth		21. Baltoro	
16. Glaciers 4, 5, 6		22. Panmah	
17. Glacier 3		23. Biafo	
18. Gangotri		24. Kero Lungma	
Lahaul-Spiti - Region C		25. Gurpaltiq	
1. Bara Shigri		26. Garumhar	
2. Sonapani		27. Kunyang	
3. Gangtang		28. Hispar	
Kolahoi - Region D		Rakaposhi-Haramosh - Region G	
1. Machoi		1. Niaro	
2. Kolahoi		2. Chogo Lungma	
Nanga Parbat - Region E		3. Timpur	
1. Sachen		4. Kotiah	
2. Chungphar		5. Mani	
3. Bazhin		6. Yengutsa	
4. Shaigiri		7. Hopar (Barpu)	
5. Toshain		8. Hinarche	
6. Rakhiot		9. Minapin	
		Ratura Mustagh - Region H	
		1. Ghulkin (Sasainil)	
		2. Pasu	
		3. Ratura	
		4. Hasanabad	
		5. Baltar	
		6. Kukoor	
		7. Karamhar	
		8. Daintar Valley Glaciers	
		9. Bad Swat Valley Glaciers	
		10. Bhort Valley Glacier	
		North Side Karakoram - Region I	
		1. Phurduka	
		2. South Shukpakunchang	
		3. North Shukpakunchang	
		4. Nameless Glacier	
		5. Sultan Chhusku	
		6. East Chamsen	
		7. West Chamsen	
		8. Aktash	
		9. Kichik Kumdan	
		10. Chong Kumdan	
		11. South Rimo	
		12. Central Rimo	
		13. Yarkand Rimo	
		14. Lungmoche-A	
		15. Lungmoche-B	
		16. Kyagar	
		17. Singye	
		18. Staghar	
		19. Urdok	
		20. Gasherbrum	
		21. Virjerab	
		22. Khurdopin	
		23. Yazghil	
		24. Malangutti Yaz	
		25. Momhil Yaz	
		26. Lupghar Yaz	
		27. Ghutulji Yaz	
		Khunjerab-Ghujerab - Region J	
		1. Parpik	
		2. Kuxsel	
		3. North Maidur	
		4. South Maidur	



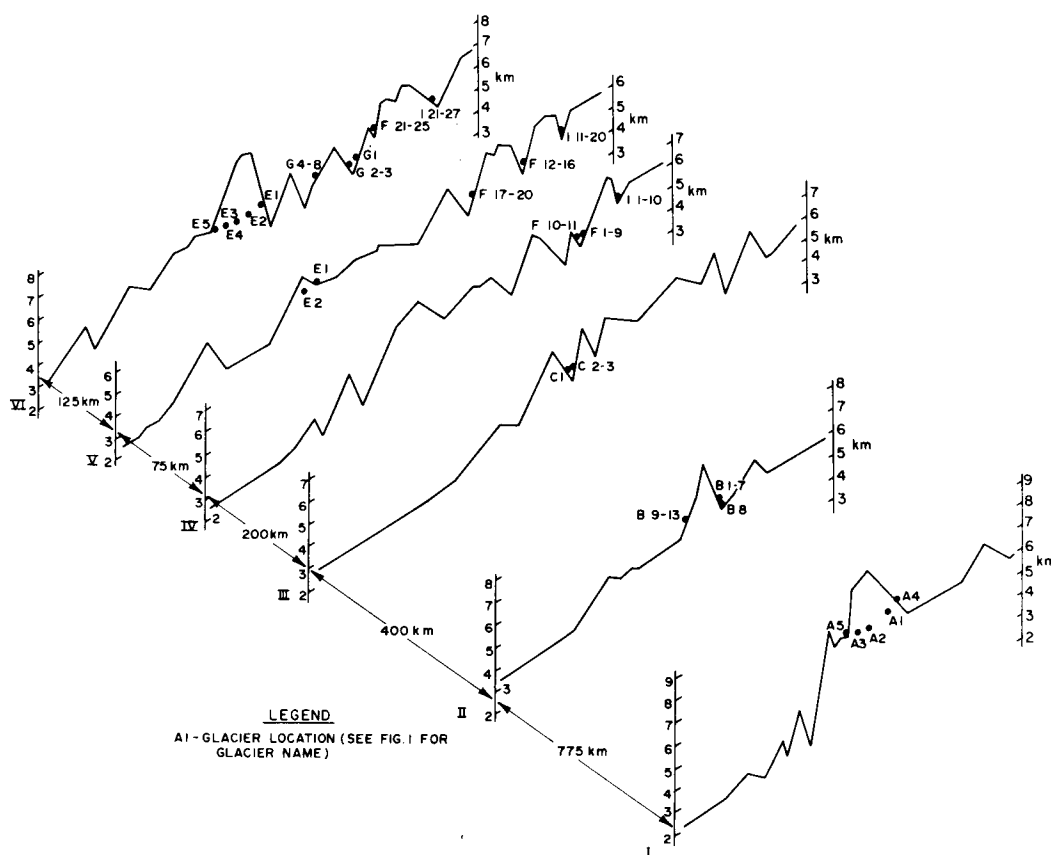


FIGURE 2. Cross sections displaying the altitudinal range and orientation of selected glaciers. Only portions of the cross sections above 2 km in elevation are shown.

notes, and photographs from his own and others' visits, documented the fluctuations of 34 glaciers in the Trans-Himalayas (Figure 1, F to J). Physical descriptions of all of these glaciers are included in his paper. Mason divided the glaciers in this area into two types—longitudinal and transverse—on the basis of their physical characteristics. Longitudinal glaciers flow generally in the wider east-west trending valleys of the range. These glaciers are relatively long and have relatively gentle surface slopes compared to the transverse glaciers. The transverse glaciers with their shorter length and steeper surface slopes flow perpendicular to the longitudinal type and therefore nearly perpendicular to the main axis of the range. Mason (1930a) summarized the fluctuation records of 21 of these glaciers and noted that several of the records displayed recognizable variations. These

variations were categorized as secular (long-term), periodic (short-term), seasonal, and accidental.

Visser and Visser-Hooft (1938, in Washburn, 1939) discussed fluctuations of 72 glaciers from the Trans-Himalayan region (Figure 1, F to J). Most of their information is drawn from Mason (1926), Mason (1930a), or their own observations during expeditions to the Karakoram in the years AD 1922, 1925, 1929-1930, and 1935. Visser and Visser-Hooft plotted glacier fluctuations as a function of time for 61 glaciers, 32 of which had previously appeared in Mason (1930a). Their study also suggested that in the period AD 1900 to 1910 most of the glaciers for which records were available were advancing, whereas in the period AD 1910 to 1920 the majority were retreating.

Mercer (1963) discussed the fluctuations of

50 glaciers: 43 from the Trans-Himalayas (Figure 1, F to J) and 7 from the Nanga Parbat Massif (Figure 1, E). Thirty-one of these glaciers appeared in either Mason (1930a) or Visser and Visser-Hooft (1938). Mercer grouped them into three classes of advance: steady, cyclic or repeated, and catastrophic. Steady advances affected primarily the longer glaciers, especially from the early 19th century until approximately AD 1920, while the cyclic or repeated advances affected primarily shorter glaciers. The catastrophic advances documented are believed primarily to have been caused by earthquakes.

## LOCAL SYNTHESSES

The Himalayan and Trans-Himalayan records of glacier fluctuation used in our study (Figures 3-6) are arranged geographically according to the groupings delineated on Figure 1. The records are displayed graphically in a fashion similar to the formats used by Mason (1930a) and Visser and Visser-Hooft (1938). Advances and retreats were determined from records of the movement of glacier termini. Due to a general absence and, in some cases, the unreliability of data concerning ice volume changes, such as downwasting and lateral spreading, the termini movements are used as the sole indication of advance or retreat. However, where ice-volume changes were interpretable, descriptions appear on the graphs. An example is Khumbu Glacier, in the Everest-Kanchenjunga area (Figure 4) whose snout remained stationary from about AD 1930 to 1956 although the glacier thinned approximately 70 m under a thick debris cover during this period (Müller, 1958).

Summaries of the fluctuation records rely upon differing numbers of observations (plotted on Figures 3-6 as data points). Wherever possible interpolation between data points was made based upon conclusions drawn from published material. Conjectural connections between data points were so labelled based upon published material or, in a limited number of cases, our own interpretation of the data. In some cases single points, for example Mrigthuni Glacier in the Garwhal (Figure 3), have lines drawn through them indicating that the data point is part of a trend in advance or retreat. Some glaciers

Tewari (1971) tabulated fluctuation data for 17 glaciers in the Himalayas and Trans-Himalayas and concluded that glaciers in the Himalayas, in general, were retreating while certain glaciers in the Trans-Himalayas displayed unexplained cyclic fluctuations.

The Himalayan and Trans-Himalayan fluctuations synthesized in this paper are presented as a reevaluation of the existing literature. Our work updates previous summaries and introduces an additional 27 glaciers for a total of 112 records of glacier fluctuations. Additional pertinent information concerning each glacier appears in the Appendix.

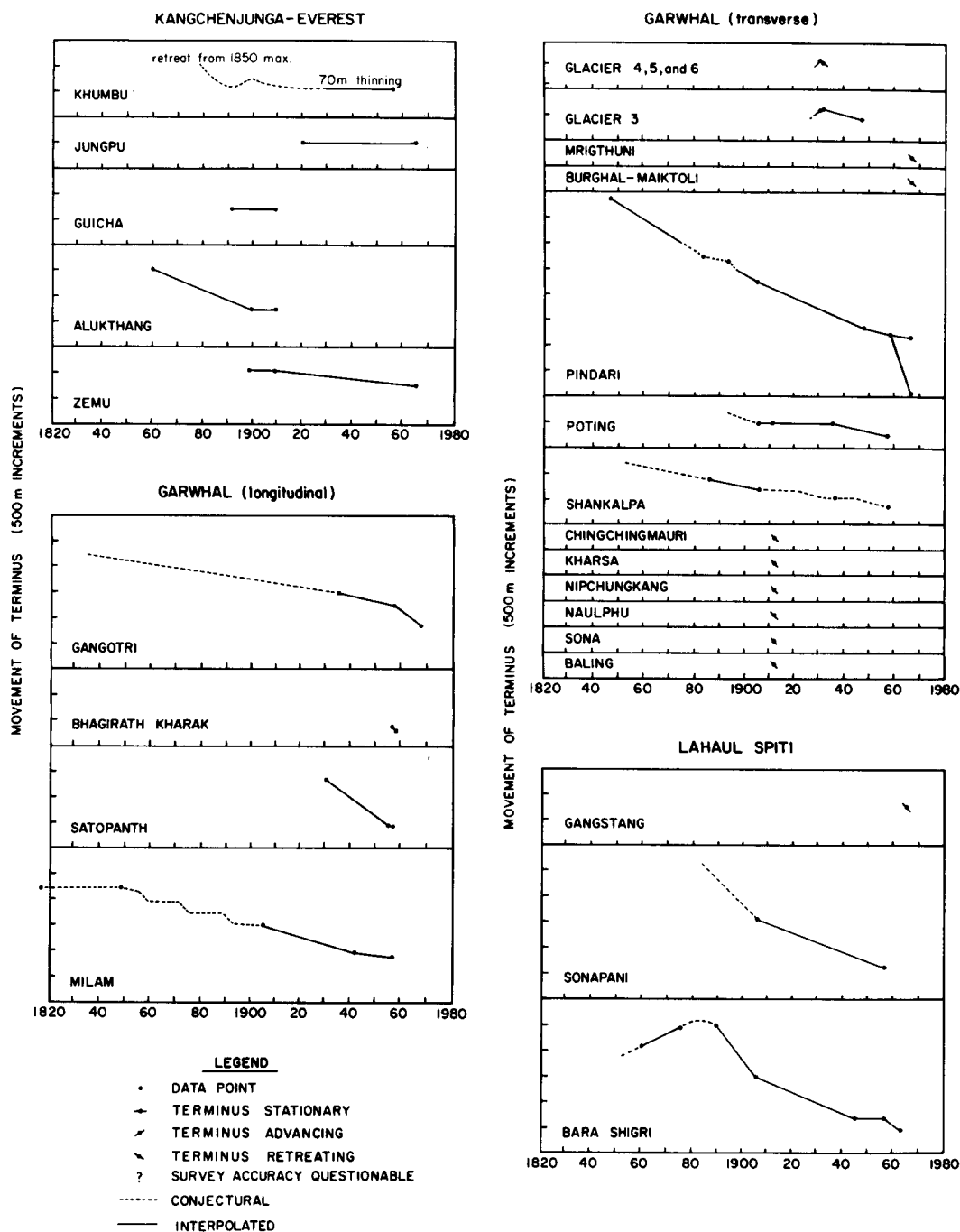
contain several data points from which more complex records can be interpreted, for example Pindari Glacier in the Garwhal (Figure 3). However, even in these cases short-term advances or retreats may have been lost because of the spacing of the recording periods. Summaries of the fluctuation records for each of the 10 geographic areas in Figure 1 are presented below.

### KANCHENJUNGA-EVEREST

Kanchenjunga-Everest glaciers (Figure 3) are commonly characterized by steep gradients, such as the vertical drop from 8500 to 4900 m over the 20-km length of Khumbu Glacier. No Everest-Kanchenjunga glaciers extend below 4000 m. Terminal portions of many of the glaciers in this region are heavily laden with debris. The relative stability of the snout positions held by Khumbu Glacier (Miller, 1965), Zemu Glacier (Bose et al., 1971), and Jungpu Glacier (Scientia, 1974), despite their recent thinning, has been ascribed to the insulating effect of this debris cover. This insulating effect is apparent even though each of these glaciers faces a different direction: Khumbu Glacier faces south, Zemu Glacier faces east, and Jungpu Glacier faces north.

### GARWHAL

The Garwhal (Figure 3) contains some of the longest glaciers in the Himalayas; an example is Gangotri Glacier which is 30 km long. Several of the glaciers in this area descend to as low as 3600 m. Garwhal glaciers for which records are available can be divided



into longitudinal and transverse types. Although the data are sparse it appears that glaciers in the Garwhal, regardless of type, have been in a state of retreat since AD 1850.

Pindari Glacier has undergone the greatest retreat of any glacier in the Himalayas. Total retreat of Pindari Glacier since AD 1850 has been 2600 m. By AD 1958 retreat of this glacier caused it to split into two tributaries.

Milam Glacier is 20 km long and is the second longest glacier in the central Himalayas. This longitudinal glacier retreated 800 m from AD 1849 to 1906 (Cotter and Brown, 1907) and a further 620 m from AD 1906 to 1957 (Jangpangi, 1958). The lower 7 km is currently debris covered. The ridged lateral moraines and a series of end moraines indicate that retreat has not been steady.

#### LAHAUL-SPITI

Lahaul-Spiti glaciers (Figure 3) are located in the monsoon-arid transition zone (Krenek and Bhawan, 1945). Therefore, they may be considered to be potentially sensitive indicators of the northern limits and intensity of the monsoon. They are commonly covered by debris as indicated by the name Bara Shigri Glacier which means debris covered in the Spiti dialect (Dutt, 1961). Since their most recent maximum, approximately AD 1880 to 1890, the three glaciers of Lahaul-Spiti for which records exist have been receding.

Bara Shigri Glacier is a transverse glacier more than 10 km in length. This glacier advanced across the valley of the Chandra River damming a lake from AD 1860 to 1893 (Egerton, 1864). Retreat rates since this last maximum have been  $62.5 \text{ m yr}^{-1}$  from AD 1890 to 1906,  $20.5 \text{ m yr}^{-1}$  from AD 1906 to 1945, and up to  $28 \text{ m yr}^{-1}$  from AD 1956 to 1963 (Walker and Pascoe, 1907; Krenek and Bhawan, 1945; Srikantia and Padhi, 1963).

#### KOLAHOI

Fluctuation records are available for only two glaciers in the Kolahoi. Of these two records the Kolahoi Glacier record is the longest. This glacier retreated 800 m from AD 1857 to 1909 and 800 m from AD 1912 to 1961 (Odell, 1963). A shorter record is available for Machoi Glacier. It advanced slightly around AD 1900 and then retreated 457 m from AD 1906 to 1957 (Tewari, 1971).

#### NANGA PARBAT

The Nanga Parbat massif is characterized by extremely steep topography, vertical gradients are on the order of  $300 \text{ m km}^{-1}$ . Glaciers in this massif (Figure 4), such as Chungphar Glacier, descend to elevations as low as 2800 m.

Glaciers of Nanga Parbat have experienced a general, but rather minor, retreat (average  $6 \text{ m yr}^{-1}$ ) since AD 1850. The only significant departures from this low retreat rate exist for Rakhiot and Chungphar glaciers which between AD 1930 and 1950 had retreat rates as high as  $30 \text{ m yr}^{-1}$ . Despite this increase in retreat rate Rakhiot Glacier (Paffen et al., 1956) and Chungphar Glacier (Finsterwalder and Pillewizer, 1939) have been characterized by velocities that within the period AD 1930 to 1950 have been as much as two or more times pre-AD 1930 velocities.

#### SOUTH SIDE OF KARAKORAM

The south side of the Karakoram (Figures 4 and 5) contains some of the longest glaciers outside of the polar regions such as Siachen (75 km long) and Hispar (61 km long). Records from this area are available for both longitudinal and transverse glaciers although records for the former are far more complete. Several types of records are displayed by glaciers in this area (Mason, 1930a): (1) general recession since AD 1880 (Hispar) and even since AD 1850 (Baltoro), (2) fluctuation between advance and retreat in the period AD 1890 to 1930 (Biafo), and (3) catastrophic movements (Garumbar).

Hispar is a longitudinal glacier that forms the main trunk glacier for the local area. The névé field that feeds this glacier also feeds Biafo Glacier and several tributary glaciers. Documentation of the movement of the terminus of Hispar Glacier, summarized by Mason (1930a), suggests that this main trunk glacier has undergone a period of long-term recession which began no later than AD 1880 and, although decelerating in rate by approximately AD 1910, is probably still continuing. Tributaries of this glacier display fluctuations of a shorter period that are, in fact, even "out of time" (Mason, 1930a: 227) with each other.

The greatest number and magnitude of fluctuations recorded for any glacier on the



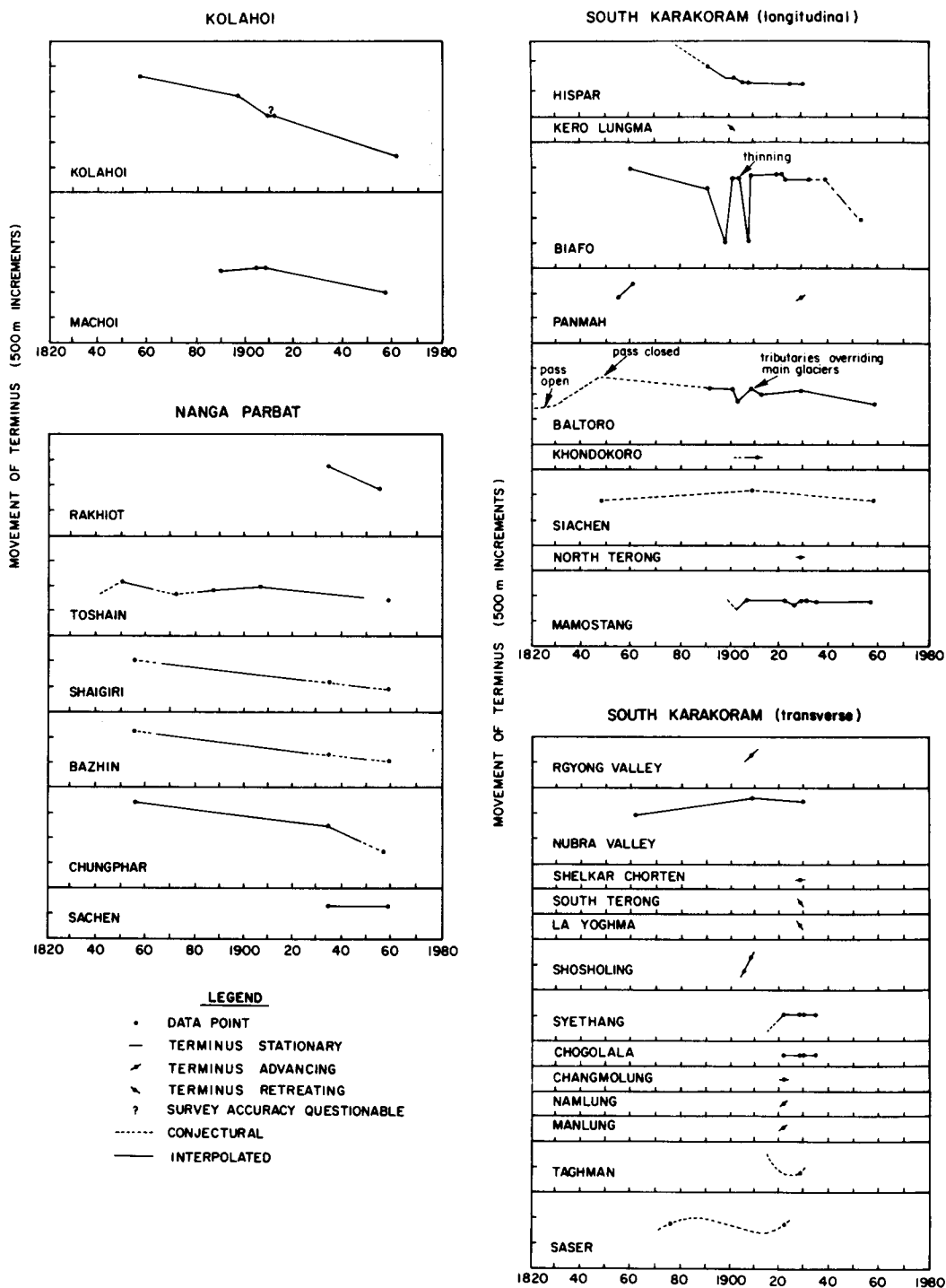


FIGURE 4. Kolahoi, Nanga Parbat, and South Karakoram glacier fluctuations.

south side of the Karakoram are displayed by Biafo Glacier. Mason (1930a) suggests that fluctuation of this longitudinal glacier may be explained by the physical controls imposed by the glacier's drainage dimensions. It has a large (64 km<sup>2</sup>) snow basin that feeds a 4-km-wide valley that tapers within 10 km of the snout to a 1.5-km-wide valley. Mason (1930a) cautions that forward advance of Biafo Glacier could cause blockage of meltwater from Baltoro and other glaciers nearby.

The south side of the Karakoram may contain several surging glaciers, such as Garumbar, which may have experienced a 2.5-km advance sometime between AD 1892 and 1925 (Mason, 1930a).

#### RAKAPOSHI-HARAMOSH

Rakaposhi-Haramosh glaciers (Figure 5) are divisible into longitudinal and transverse types. Glaciers in the area descend to an altitude of 2000 m and have been fluctuating between retreat and advance since at least AD 1850. In general, the AD 1960 margins of these glaciers are quite close to the AD 1850 positions. However, the 50 km long longitudinal Chogo Lungma Glacier reached its greatest post-AD 1300 maximum soon after AD 1912 (Kick, 1962).

Hewitt (1969) suggests that many of the glaciers of Rakaposhi-Haramosh have surged. Yengutz Glacier surged 3.2 km in 8 d during AD 1902/03 (Hayden, 1907), and Kutiah Glacier surged 12 km in 2 mon during AD 1953 (Desio, 1954). Minapin Glacier may also be a surging glacier (Mason, 1935).

#### BATURA MUSTAGH

Records for both longitudinal and transverse glaciers are available from the Batura Mustagh area. Fluctuations of Batura Mustagh glaciers (Figures 5 and 6) fall into two categories. The first includes those glaciers in the Baltar and Kukuar valleys. These valleys were free of ice in the early 19th century and by the late 19th century contained glaciers at their maximum post-AD 1820 extents. The second category includes glaciers

in Daintar Valley and Pasu, Ghulkin, and Batura glaciers which reached their maximum during the early part of the 20th century.

Batura Mustagh contains some glaciers that have been known to surge, such as Hasanabad Glacier, which advanced 9.7 km in 2.5 mon during AD 1903 (Hayden, 1907). By AD 1954, Hasanabad Glacier had broken into several tributaries in response to the total 7-km retreat it experienced from AD 1892 to 1954 (Paffen et al., 1956).

#### NORTH SIDE OF KARAKORAM

Records for both longitudinal and transverse glaciers are available from the north side of the Karakoram (Figure 6). In general, records of glacier fluctuations for this area are short, except for Chong Kumdan, Aktash, and Kichik Kumdan glaciers. Mason (1935) summarized the available data for these glaciers and concluded that periodic variations of 90 yr (Chong Kumdan), 55 yr (Aktash), and 35 yr (Kichik Kumdan) within the time period AD 1830 to 1940 were differentiable for these glaciers. Tewari (1971) suggested cycles within the period AD 1812 to 1958 of 51 yr (Chong Kumdan), 55 yr (Aktash), and 45 yr (Kichik Kumdan). Mason (1930a) suggests that periodicity in the fluctuation record of Chong Kumdan Glacier, in particular, is due to its large snow basin reservoir which is capable, by means of outflow to the glacier, of exaggerating changes in snow input. All three glaciers have blocked the course of the Shyok River at several times, impounding large lakes (Mason, 1929). Subsequent bursting of these dams has caused a grave threat to the inhabitants below. In addition, Kichik Kumdan is believed to have surged 2.4 km in 7 mon during AD 1935/36 (Mason, 1940a).

#### KHUNJERAB-GHUJERAB

Glaciers of Khunjerab-Ghujerab (Figure 6) do not descend below 4000 m. Records for the glaciers in this area only exist for AD 1925 (Mason, 1930a) and thus no interpretation can be made.

### REGIONAL SYNTHESIS

Trends in the history of Himalayan and Trans-Himalayan glacier fluctuations are revealed in this study by reviewing the records

of glacier fluctuations previously summarized as (1) a composite (112 glaciers), (2) groupings differentiated by glacier type (27 longi-

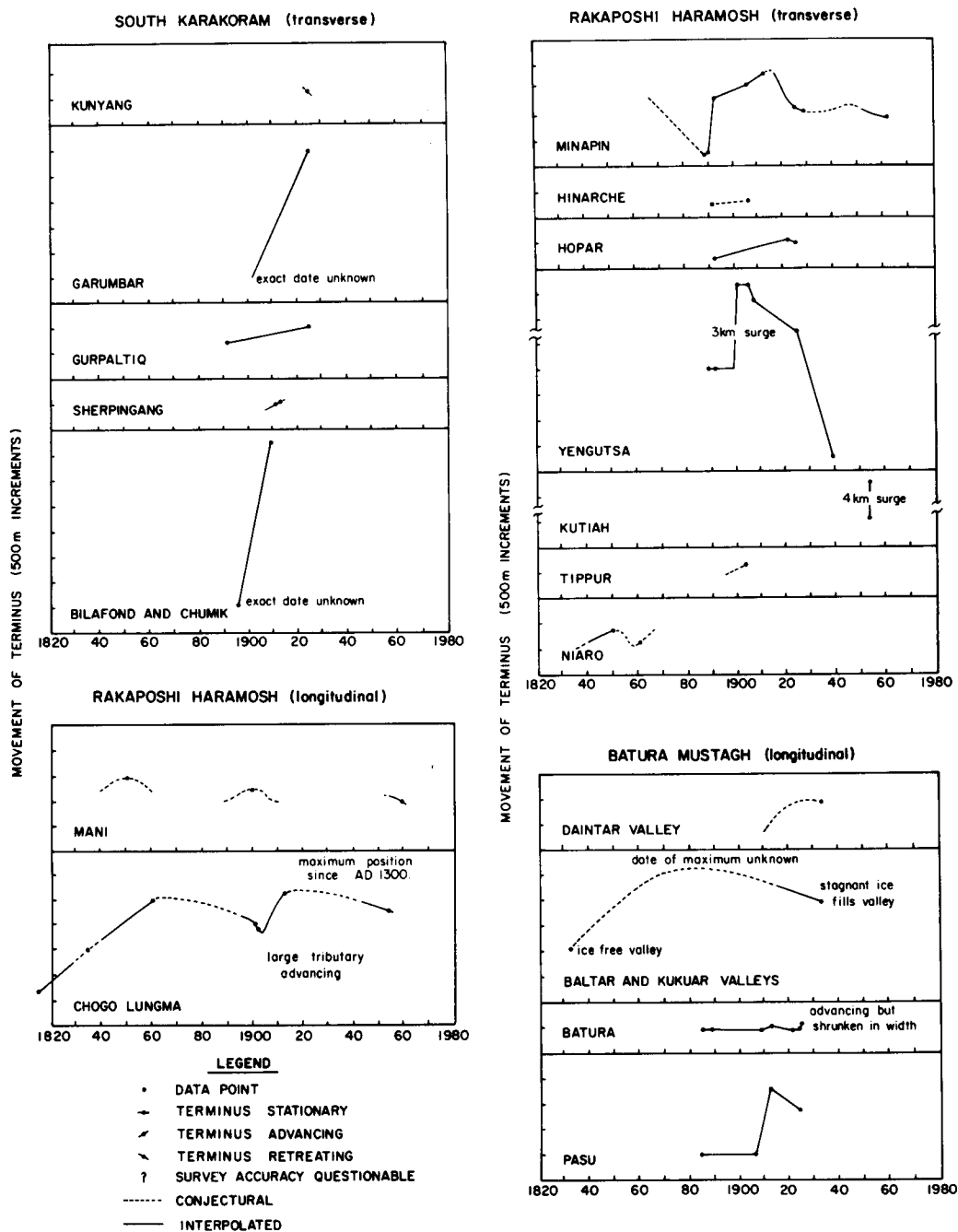


FIGURE 5. South Karakoram, Rakaposhi-Haramosh, and Batura Mustagh glacier fluctuations.

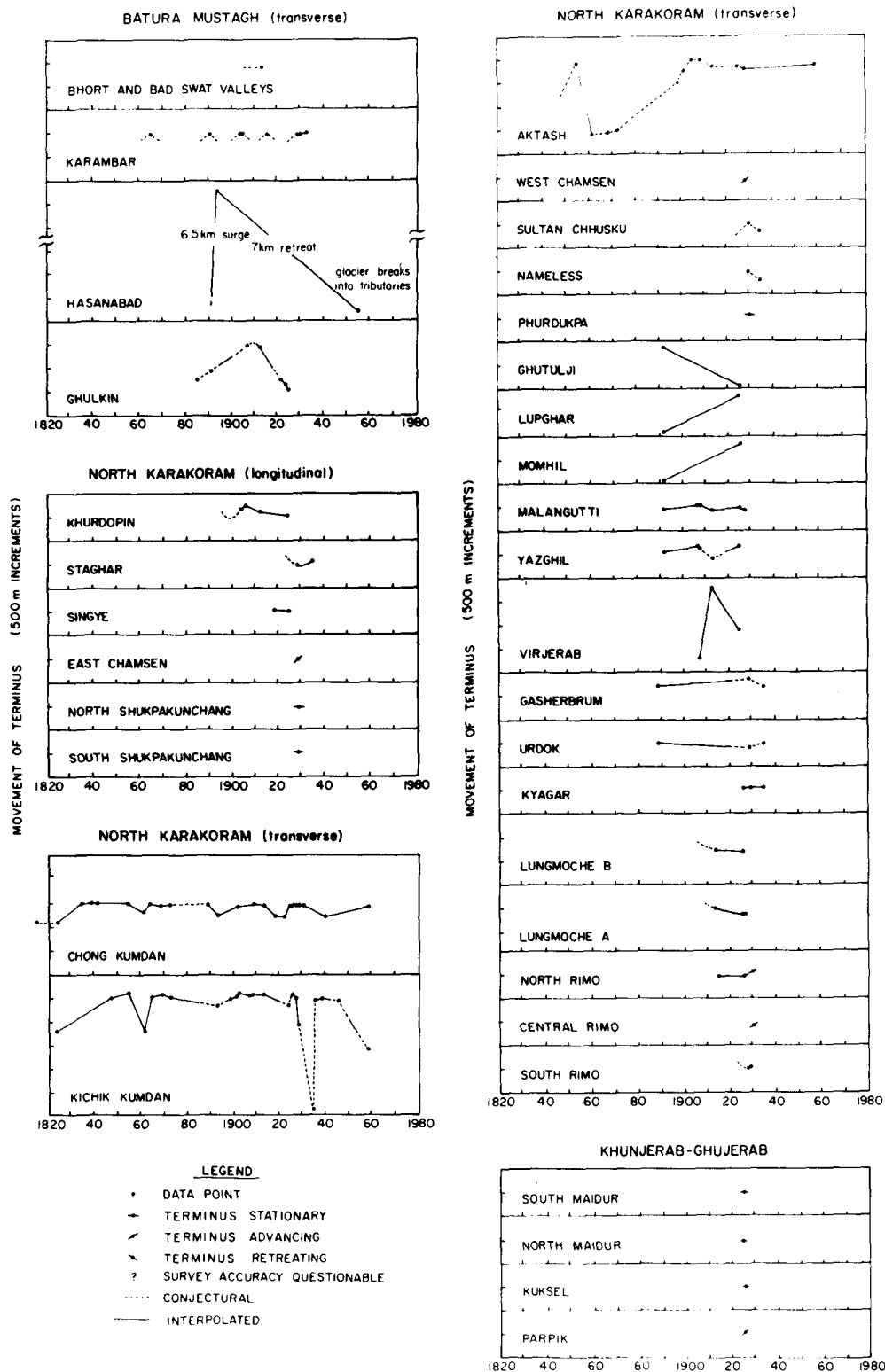


FIGURE 6. Batura Mustagh, North Karakoram, and Khunjerab-Ghujerab glacier fluctuations.

tudinal and 68 transverse glaciers), and (3) groupings differentiated by regional setting (34 Himalayan and 74 Trans-Himalayan glaciers). Data used in this synthesis are presented in the form of frequency diagrams of glacier activity (Figure 7a-e) in which glacier activity is plotted for 10-yr periods. The 10-yr glacier activity trends fall into three cate-

gories: advance, retreat, and standstill, each referring to the relative position of the terminus of the glacier. The data are presented as percentages to allow absolute comparison of 10-yr periods despite differing numbers of observations per 10-yr period. The total numbers of observations per 10-yr period are also included in Figure 7 to prevent misinterpreta-

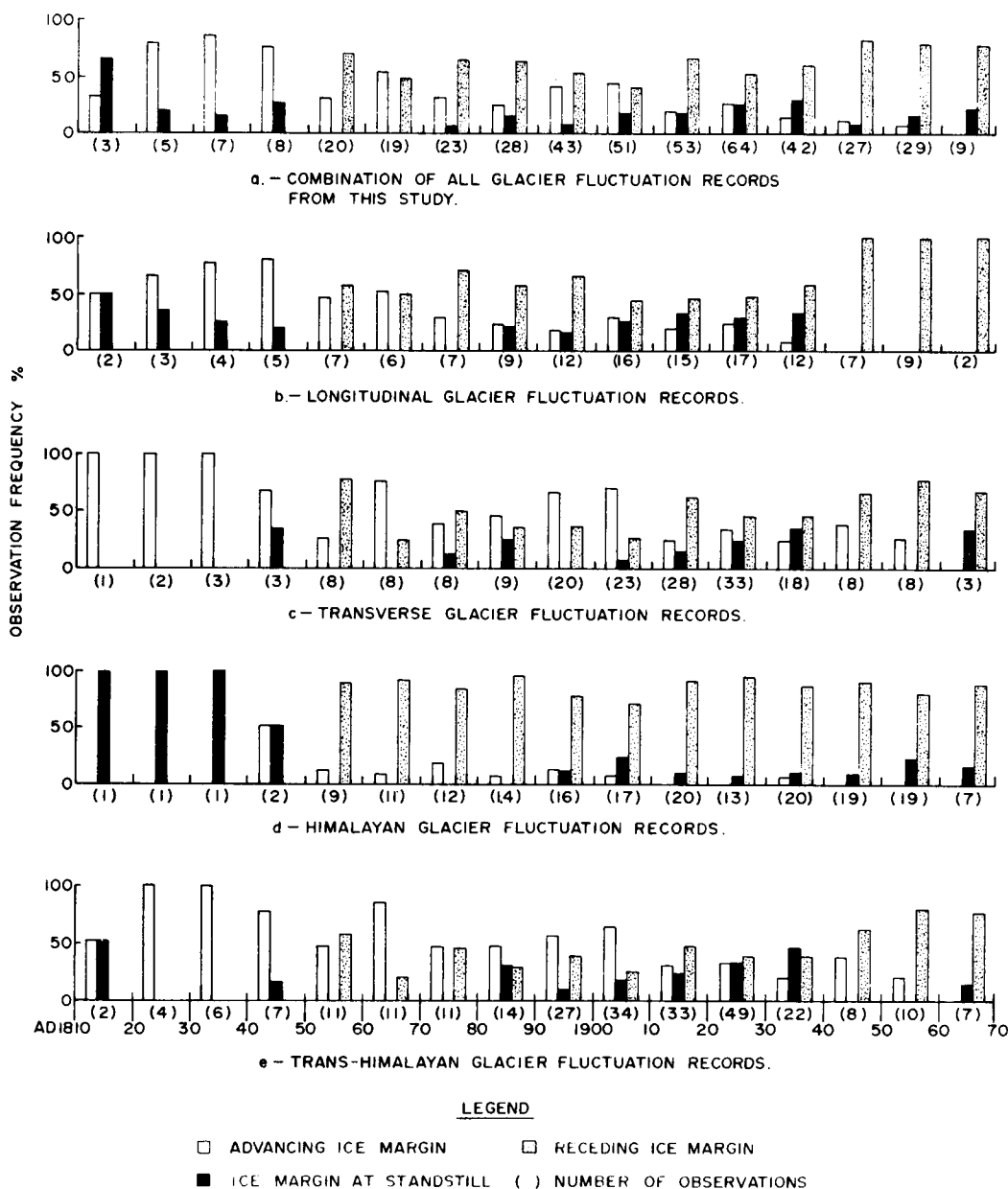


FIGURE 7. Glacier activity frequencies.

tion of the percent frequency diagrams introduced by different sample populations per 10-yr period. In general, glacier observations have increased since AD 1850 with the peak in the observational period occurring from AD 1910 to 1930. Records of surging glaciers have been excluded from Figure 7 because these glacier movements may be spurious and represent only a small fraction of the fluctuation record. In addition, records from glaciers prior to AD 1850 are also not included in this synthesis due to the sparsity of observations from AD 1810 to 1850. Results from Figure 7a-e follow.

Review of the composite record of all glacier fluctuations (Figure 7a) suggests that as a regional grouping the dominant regime since AD 1850 has been retreat. However, the period AD 1870 to 1970 is characterized by several secondary trends in advance and standstill. In addition, when the number of advancing and stationary glacier termini are combined they exceed the number of retreating glacier termini during the period AD 1920 to 1940, suggesting a near reversal in the dominance of the retreat regime for this period. Although review of the composite record yields a general impression of glacier regimes as a function of time, details in the record may be lost.

Analyses of the fluctuation of glaciers of different types and/or regional setting provide greater detail than that available from the composite record. A comparison of the records for longitudinal (Figure 7b) and transverse (Figure 7c) glaciers demonstrates that these physically different types of glaciers have different fluctuation histories. Longitudinal glaciers have been in a general state of retreat that began approximately AD 1850. However, if advance and standstill records for longitudinal are combined, they dominate from AD 1900 to 1930 and provide a strong secondary regime from AD 1880 to 1900 and AD 1930 to 1940. Transverse glaciers retreated as a group from AD 1850 to 1870. Combined advance and standstill dominated transverse glaciers from AD 1870 to 1910 and from AD 1920 to 1930; retreat characterized

the periods AD 1910 to 1920 and AD 1940 to 1960, and it probably continues today. Causes for the differences in the records of the two types of glaciers are only conjectural at this time, but we suggest that the apparently more complex record displayed by the transverse glacier group may be due to the fact that these glaciers are shorter, flow perpendicular to oncoming atmospheric circulation patterns, and have steeper gradients.

Fluctuation records of glaciers in the Himalayas (Figure 7d) differ from those of glaciers in the Trans-Himalayas (Figure 7e). Most glaciers in the Himalayas (Kanchenjunga-Everest, Lahaul-Spiti, Garwhal, Kolahoi, and Nanga Parbat) have been in a general state of retreat since AD 1850. Those in the Trans-Himalayan grouping (south side of the Karakoram, Rakaposhi-Haramosh, Batura Mustagh, and north side of the Karakoram) were either in retreat or advance from AD 1850 to 1880, reflected near-equivalent influences of retreat, standstill, and advance regimes from AD 1880 to 1940 (combined standstill and advance records dominated), and have retreated since AD 1940. Tewari (1971), sampling a much smaller number of glaciers than the number sampled in this study, found similarly that the Himalayan glaciers were in a general state of retreat while the Trans-Himalayan glaciers have a more complicated record.

Glacier classifications other than the longitudinal and transverse types of Mason (1930a) have been suggested in the Himalayas and Trans-Himalayas. Visser and Visser-Hooft (1938, in Washburn, 1939) subdivided Karakoram glaciers into classes using the characteristics of their nourishment area (basin-reservoir, plateau-reservoir, incised-reservoir, and avalanche types), whereas Pillewizer (1958) subdivided Himalayan and Karakoram glaciers into classes on the basis of their type of movement (block and flow types). Comparison of records of glacier fluctuations implementing these and other types of classifications may prove to be as informative as the longitudinal/transverse and the Himalayan/Trans-Himalayan comparison.

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## APPENDIX

### *Glacier Data*

Glacier name	Location index (see Figure 1)	Type	Length (km)	Exposure	Snout elev. (m)	Last date examined	References
Everest-Kanchenjunga (Region A)							
Zemu	1	—	25	E	4000	1965	9,28,29,50, 66,79,96
Alukthang	2	—	—	S	4000	1909	6,28,29,66,79
Guicha	3	—	—	SE	—	1909	66,79
Jungpu	4	—	—	N	> 4000	1968	102
Khumbu	5	—	20	S	4900	1963	8,57,74,79,80, 81,82,84,125
Garwhal (Region B)							
Baling	1	TR	5	NE	3250	1912	37,38,39
Sona	2	TR	6	NE	3380	1912	37,38,79
Naulphu	3	TR	6	NE	4000	1912	38,79
Nipchungkang	4	TR	6	NE	3860	1912	38,79
Kharsa	5	TR	6	NE	3810	1912	38,79
Chingchingmauri	6	TR	6	NE	4170	1912	38,79
Shankalpa	7	TR	13	SE	4300	1957	14,29,43,44, 54,55,79,115
Milam	8	L	20	SW	3390	1962	1,14,29,43,44, 49,54,79,108, 109,110,115
Poting	9	TR	5	SW	3650	1957	14,29,37,54, 79,115
Kaphni	10	TR	3	S	3800	1958	108,115,118
Pindari	11	TR	5	SW	3650	1966	3,10,14,29, 54,79,83,98, 108,111,115, 116,117
Burghal-Maiktoli	12	TR	4	S	—	1966	112
Mrigthuni	13	TR	8	S	—	1966	113
Bhagirath-Kharak	14	L	18	E	3650	1959	54,79,94,115
Satopanth	15	L	16	NE	3800	1959	54,79,104, 108,115
Glacier IV, V, VI	16	TR	2-3	N	4200	1932	7,31
Glacier III	17	TR	2-3	N	4200	1957	7,31,54,79,115
Gangotri	18	L	30	NW	3880	1967	5,36,54,79, 114,115,127
Lahaul-Spiti (Region C)							
Bara Shigri	1	TR	> 10	N	4080	1963	21,22,29,30,40, 64,70,79,105, 115,126,130

Glacier name	Location index (see Figure 1)	Type	Length (km)	Exposure	Snout elev. (m)	Last date examined	References
Sonapani	2	TR	15	SE	3856	1957	29,40,65,79, 115,126
Gangstang	3	TR	6	SE	4420	1963	11,79,85
Kolahoi (Region D)							
Machoi	1	TR	3	N	3733	1957	86,89,95,115
Kolahoi	2	TR	5	NW	3660	1973(?)	2,29,79,86,87
Nanga Parbat (Region E)							
Sachen	1	—	6	E	3350	1958	25,60,67,77, 79
Chungphar	2	—	8	SE	2850	1958	25,60,67,77, 79
Bazhin	3	—	7	SE	3250	1958	25,60,67,77, 79
Shaigiri	4	—	6	S-SE	3600	1958	25,60,67,77, 79
Toshain	5	—	7	SE	3700	1958	24,25,60,67, 77,79
Rakhiot	6	—	12	N	3160	1954	27,59,77,79, 92,93
South Side of the Karakoram (Region F)							
South Saser	1	TR	11	W	4850	1922	33,77,78,124
North Saser	2	TR	6	SE	4800	1922	33,77,78,124
Taghman	3	TR	6	NW	4725	1929	124
Manlung	4	TR	9	W	5000	1922	124
Namlung	5	TR	3	W	—	1922	78,124
Changmolung	6	TR	7	NW	5200	1922	78,124
Chogolala	7	TR	3	N	4850	1922	78,124
Syethang	8	TR	6	N	4900	1922	78,124
Mamostang	9	L	16	S	4400	1935	78,115,124
Shosholing	10	TR	5(?)	E	—	1909	68,77
La Yoghma	11	TR	15	NE	4000	1929	124
South Terong	12	TR	21	NW	4050	1929	78,124
Shelkar Chorten	13	TR	21	W	—	1929	78,124
North Terong	14	L	25	S	3960	1929	78,124
Major Tongue in Nubra Valley	15	—	—	—	—	1929	124
Siachen	16	L	75	SE	3705	1958	16,51,68,77, 78,115,120, 137
Rgyong Valley	17	TR	18	W	4000	1909	68,77,78,124
Bilafond and Chumik	18	TR	23	S	3800	1912	68,77,78, 124,137
Sherpingang	19	TR	15(?)	SW	—	1913	18,77,78,137
Khondokoro	20	L	18(?)	S	—	1911	77,78,137
Baltoro	21	L	58	W	3530	1958	13,17,18,27, 32,73,77,78, 91,124
Panmah	22	TR	27	S	3625	1929	19,32,73,77, 124
Biafo	23	L	59	SE	3160	1953	4,13,17,23, 32,46,62,73, 77,78,91, 124



Glacier name	Location index (see Figure 1)	Type	Length (km)	Exposure	Snout elev. (m)	Last date examined	References
Kero Lungma	24	L	19	E	3600	1902	63,124
Gurpaltiq	25	TR	4	N	3320	1925	13,121,124
Garumbar	26	TR	8	N	3230	1925	13,47,75,77, 121,124
Kunyang	27	TR	24	S	3450	1925	121,124
Hispar	28	L	62	W	3000	1930	13,29,41,73, 77,121,124, 133,135

#### Rakaposhi-Haramosh (Region G)

Niaro	1	TR	6	NE	—	1861	32,77
Chogo Lungma	2	L	50	E	2075	1954	18,32,58,60, 61,73,77,78, 88,120,124, 134,136
Tippur	3	TR	7	N	2900	1929	77,78,124, 134
Kutiah	4	TR	12	S	—	1953	20,47,77
Mani	5	L	10	NW	—	1958	77,129
Yengutsa	6	TR	11	N	3250	1939	13,29,41,47, 73,77,78, 124,133
Hopar	7	TR	21	N	2300	1925	13,47,73,77, 120
Hinarche	8	TR	6	S	—	1907	13,29,41, 77,78
Minapin	9	TR	16	N	2200	1961	12,13,29,41, 47,71,73,77, 79,99,124

#### Batura Mustagh (Region H)

Ghulkin	1	L	18	E	2440	1925	73,78,121,124
Pasu	2	L	25	E	2550	1925	73,78,121,124
Batura	3	L	58	E	2460	1925	63,73,77,78, 124
Hasanabad	4	TR	20	S	2225	1954	41,73,77,78, 90,124
Baltar	5	L	12	W	3000 ±	1933	77,78,101
Kukuar	6	L	18	S	3000 ±	1933	77,78,101
Karambar	7	TR	25	W	3000 ±	1933	48,75,77, 78,100
Daintar Valley	8	L	12	SE	3000 ±	1933	77,78,101
Bad Swat Valley	9	TR	5	W	3000 ±	1933	77,78,100
Bhort Valley	10	TR	10	W	3000 ±	1933	77,78,100

#### North Side of the Karakoram (Region I)

Purdukpa	1	TR	18	E	4900	1935	124
So. Shukpakunchang	2	L	27	SE	4640	1930	124
No. Shukpakunchang	3	L	25	E	4560	1930	124
Nameless	4	TR	7	E	4900	1935	124
Sultan Chhusku	5	TR	14	NE	4300	1935	47,77,124
East Chamsen	6	L	11	N	4700	1929	122,124

Glacier name	Location index (see Figure 1)	Type	Length (km)	Exposure	Snout elev. (m)	Last date examined	References
West Chamsen	7	TR	15	NE	4700	1929	122,124
Aktash	8	TR	8	E	4575	1958	33,34,42,43, 51,68,69,72, 73,74,75,77, 103,115,132
Kickik Kumdan	9	TR	11	E	4635	1958	26,33,34,35, 39,42,45,47, 51,68,69,73, 74,75,76,77, 78,97,103, 106,107,115, 119,124,132
Chong Kumdan	10	TR	21	E	4715	1958	18,26,33,34, 35,47,51,52, 53,56,68,69, 72,73,74,75, 78,97,103, 107,115,120, 124,132
South Rimo	11	TR	24	E	4825	1935	15,73,77,78, 124,132
Center Rimo	12	TR	40	E	5000	1935	15,73,77,78, 124,132
Rimo-Yarkand	13		20	NE	5213	1932	15,73,77,78
Lungmoche-A	14	TR	11	NW	5025	1926	73,78,124, 132
Lungmoche-B	15	TR	15	NW	5030	1926	73,78,124, 132
Kyagar	16	TR	22	N	4775	1935	19,73,78, 124
Singye	17	L-TR	24	NW	4575	1935	19,78,124
Staghar	18	L-TR	18	NW	4425	1935	19,78,124
Urdok	19	TR	23	NW	4365	1935	19,78,124, 138
Gasherbrum	20	TR	21	NE	4345	1935	19,78,124, 138
Virjerab	21	TR	40	NW	3450	1925	73,78,121, 124
Khurdopin	22	L-TR	47	NW	3250	1925	73,77,78, 121,123,124
Yazghil	23	TR	31	N	3190	1925	73,77,78, 121,124
Malangutti Yaz	24	TR	23	N	2900	1925	73,77,78, 121,124
Momhil Yaz	25	TR	32	N	2900	1925	73,121,124
Lupghar Yaz	26	TR	12	N	3050	1925	73,124
Ghutulji Yaz	27	TR	13	N	3350	1925	73,124

#### Khunjerab-Ghujerab (Region J)

Parpik	1	TR	14	W	4360	1925	73,77,78, 124
Kuksel	2	TR	7	E	4600	1925	73,124

Glacier name	Location index (see Figure 1)	Type	Length (km)	Exposure	Snout elev. (m)	Last date examined	References
North Maidur	3	TR	10	N	4550	1925	73,124
South Maidur	4	TR	3	S	4800	1925	73,124

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